Evolutionary Model For Convergent Margins Facing Large Ocean Basins: Mesozoic Baja California, Mexico

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Santa Barbara, CA

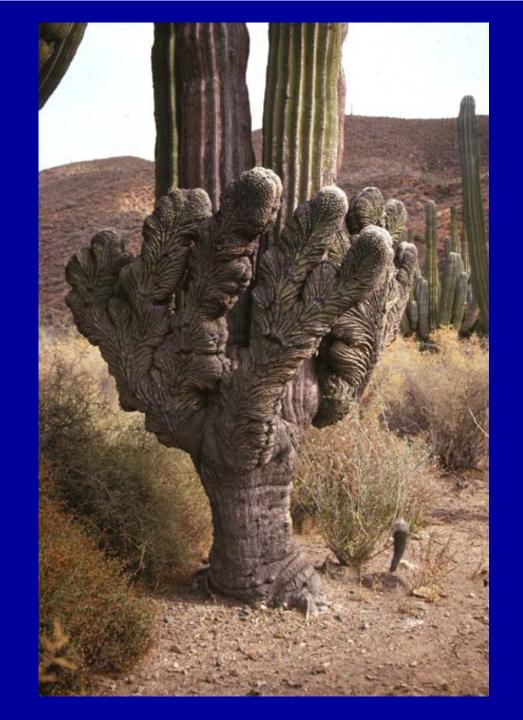


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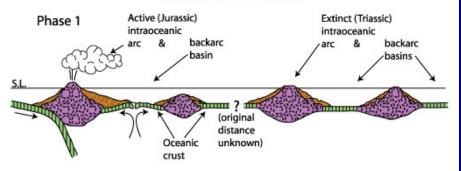
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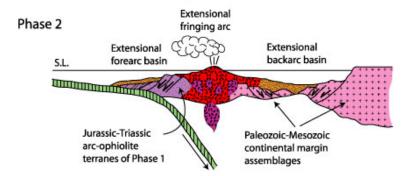




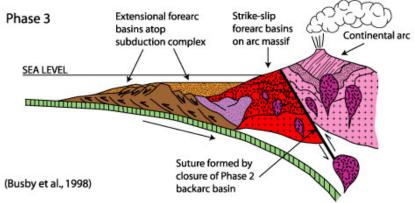
Strongly Extensional Arc

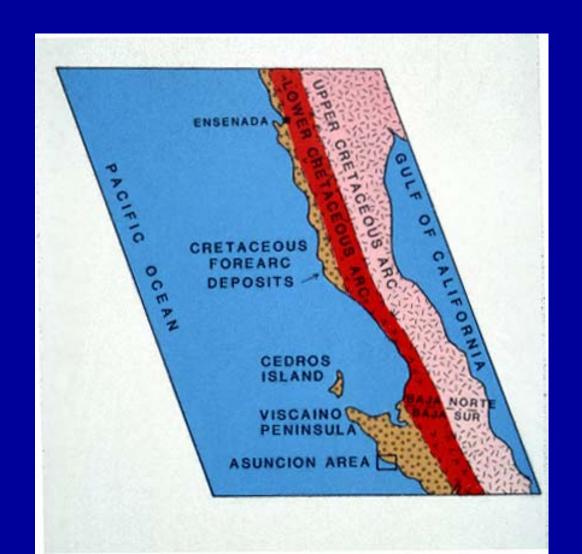


Moderately Extensional Arc

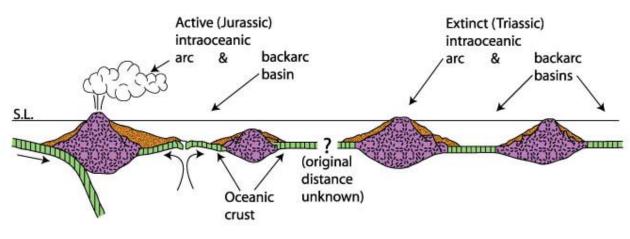


Compressional Arc





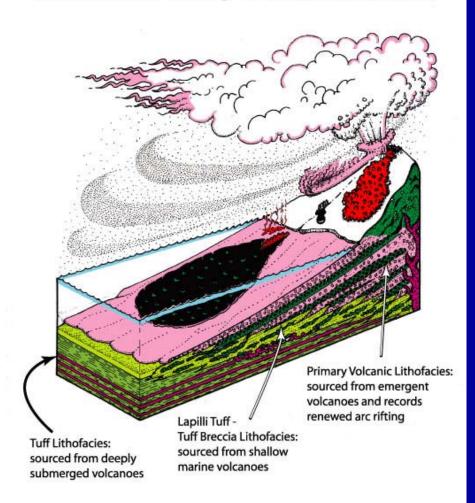
Phase 1: Strongly Extensional Arc

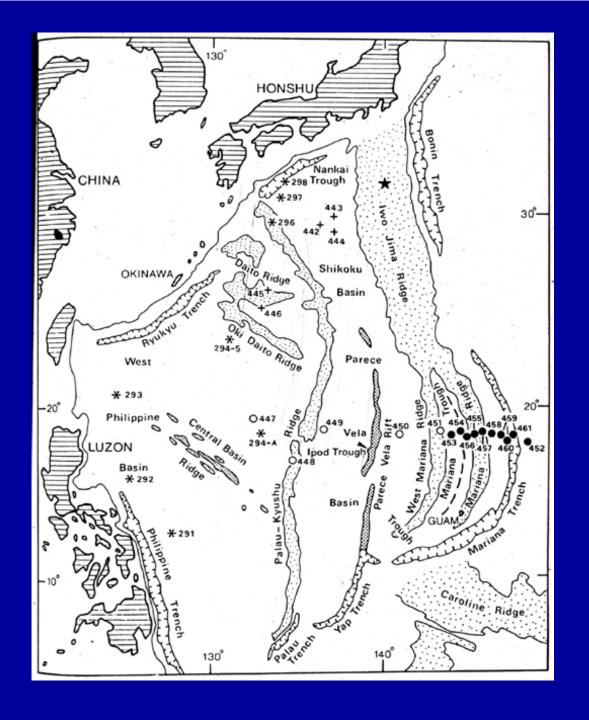


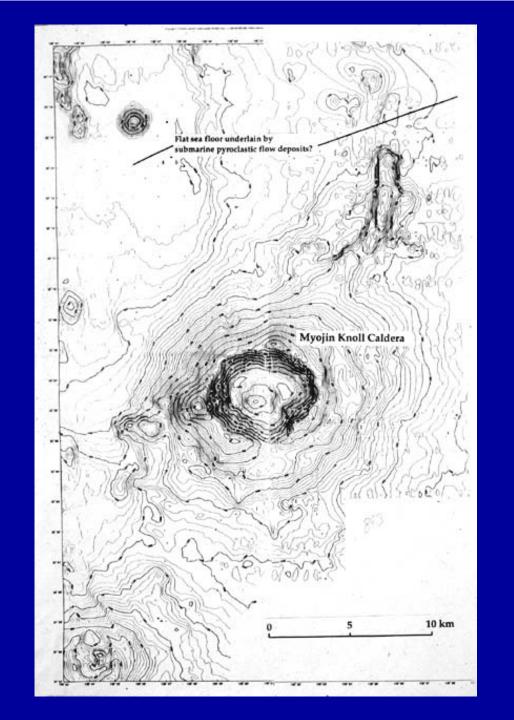
(Busby et al., 1998)

PROGRADATIONAL BACKARC APRON

Deep marine pyroclastic wedge builds from growing arc onto rifted arc crust and "steaming" oceanic crust (Busby-Spera, 1988)







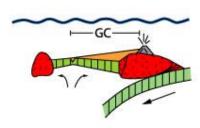
Backarc Apron Tectonic Model, Gran Canon Formation

(modified from Critelli, Marsaglia and Busby, 2002)



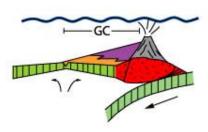
I. Rifting of oceanic arc basement

Choyal Formation



II. Nascent backarc basin (Cedros Island ophiolite, ~173 Ma) behind a deeply-submerged nascent oceanic arc

Tuff Lithofacies

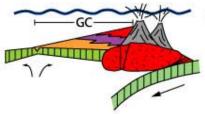


III. Growth of oceanic arc toward sea level and continued backarc spreading

Lapilli tuff - Tuff breccia Lithofacies

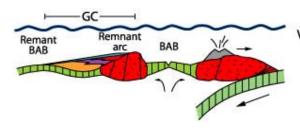
Backarc Apron Tectonic Model, Gran Canon Formation

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IV. Backarc spreading center migrates too far from subducting slab, leading to renewed arc rifting

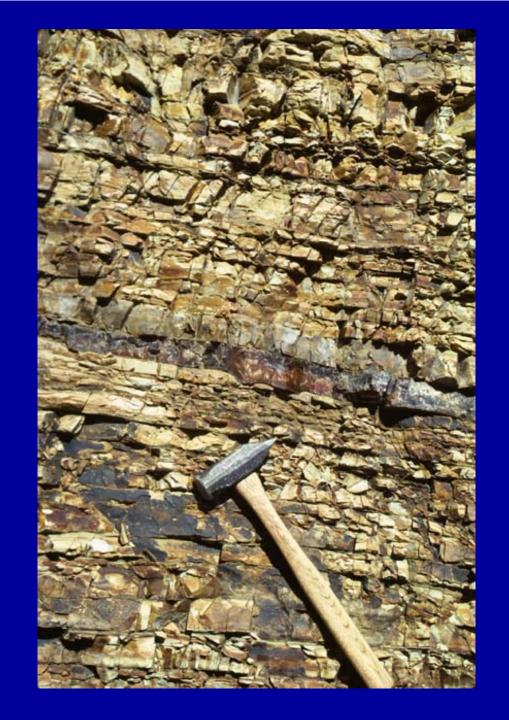
Primary Volcanic Lithofacies, ~164 Ma



V. Rifting succeeds, isolating backarc apron from active volcanoes; it is blanketed with thin cover of sand

Epiclastic Lithofacies



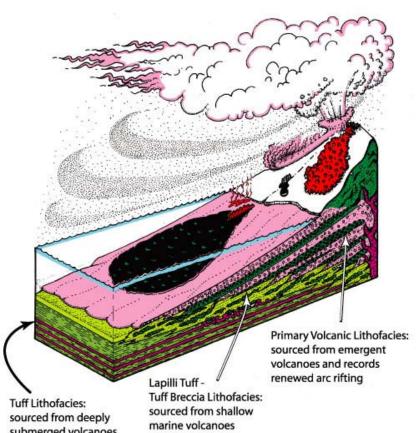






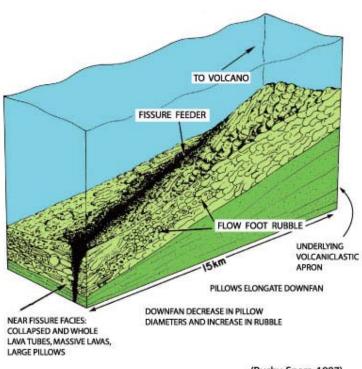
PROGRADATIONAL BACKARC APRON

Deep marine pyroclastic wedge builds from growing arc onto rifted arc crust and "steaming" oceanic crust (Busby-Spera, 1988)



submerged volcanoes

Progradational Backarc Apron: Primary Volcanic Lithofacies records renewed arc rifting



(Busby-Spera, 1987)





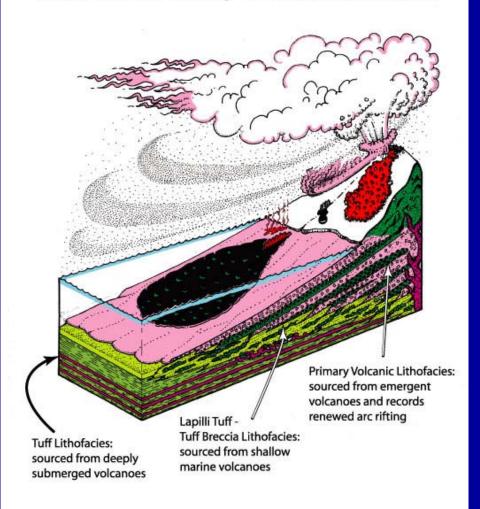




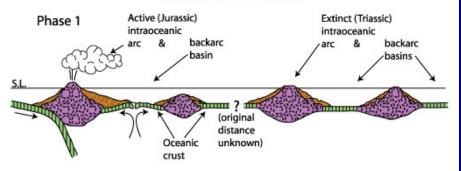


PROGRADATIONAL BACKARC APRON

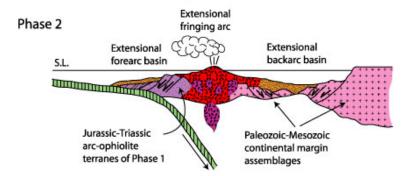
Deep marine pyroclastic wedge builds from growing arc onto rifted arc crust and "steaming" oceanic crust (Busby-Spera, 1988)



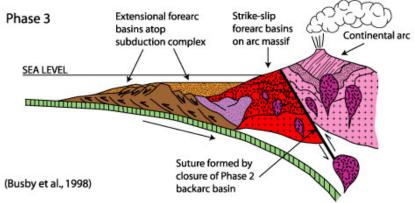
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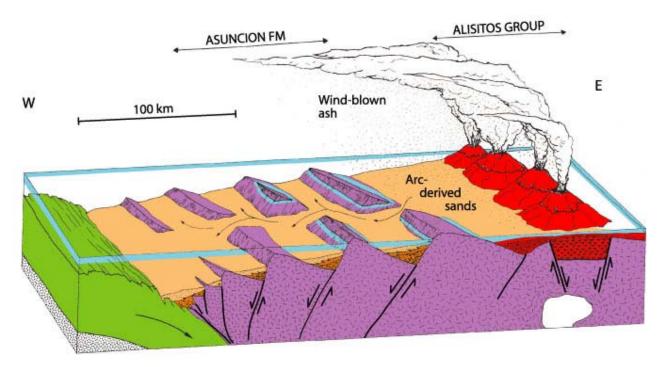


Moderately Extensional Arc



Compressional Arc





Phase 2 Extensional Arc and Extensional Forearc Basin (Busby et al., 1998)

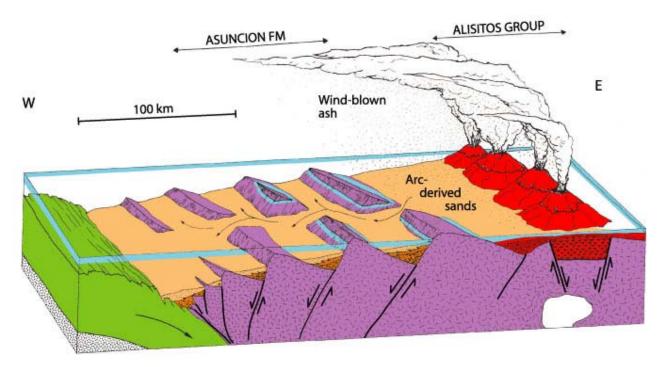
Cretaceous Alisitos arc, western Peninsular Ranges:

I. EXTENSIONAL OCEANIC ARC

Intermediate to silicic explosive volcanism, culminating in caldera-forming ignimbrite eruptions

II. RIFTED OCEANIC ARC

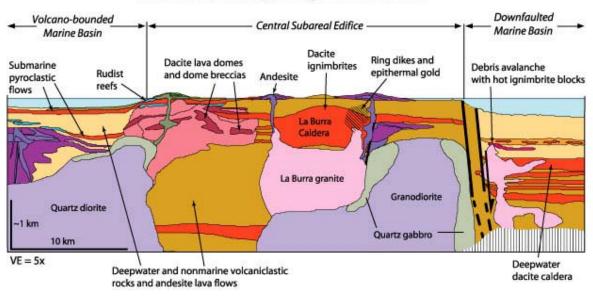
Mafic effusive and hydroclastic rocks and widespread dike swarms



Phase 2 Extensional Arc and Extensional Forearc Basin (Busby et al., 1998)

Extensional Oceanic Arc (Alisitos Group)

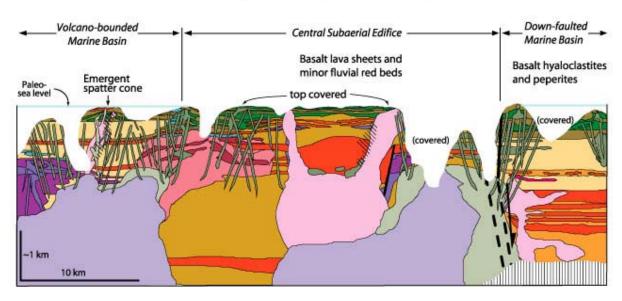
Time Slice 1: Normal faulting, high rates of subsidence (1 km/my), and ignimbrite calderas



(Adams and Busby, 1998)

Extensional Oceanic Arc (Alisitos Group)

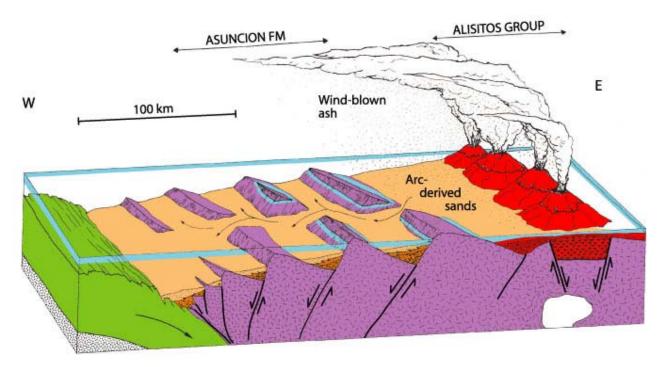
Time Slice 2: Arc rifting, mafic diking, and outpouring of basalts



True (scaled) cross-section with 5X vertical exagerration. Top covered. (Adams and Busby, 1998)

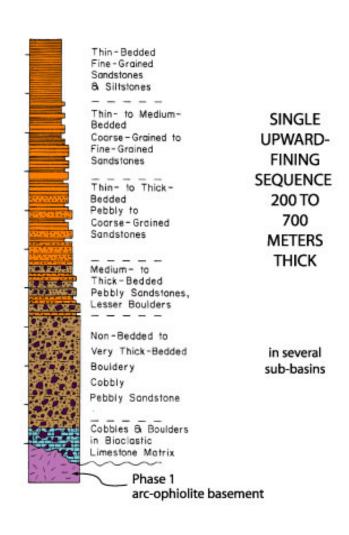


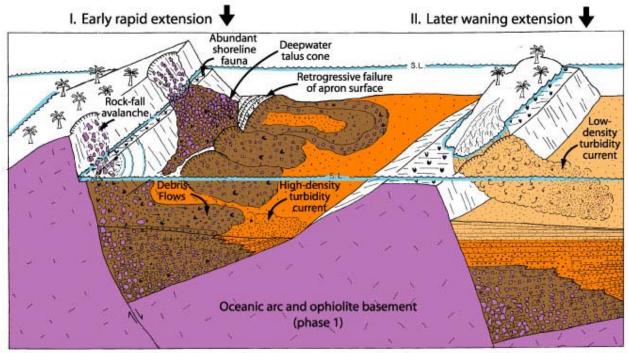




Phase 2 Extensional Arc and Extensional Forearc Basin (Busby et al., 1998)

Extensional Forearc, Phase 2 Early Cretaceous Asuncion Formation, Vizcaino Peninsula (Busby-Spera and Boles, 1986)





 Screen cones build directly from coastal fault scarps onto graben floors at bathyal water depths

II. Fine-grained turbidite wedge shed from eroded horst blocks

Extensional Forearc slope Apron Deposits of Phase Two

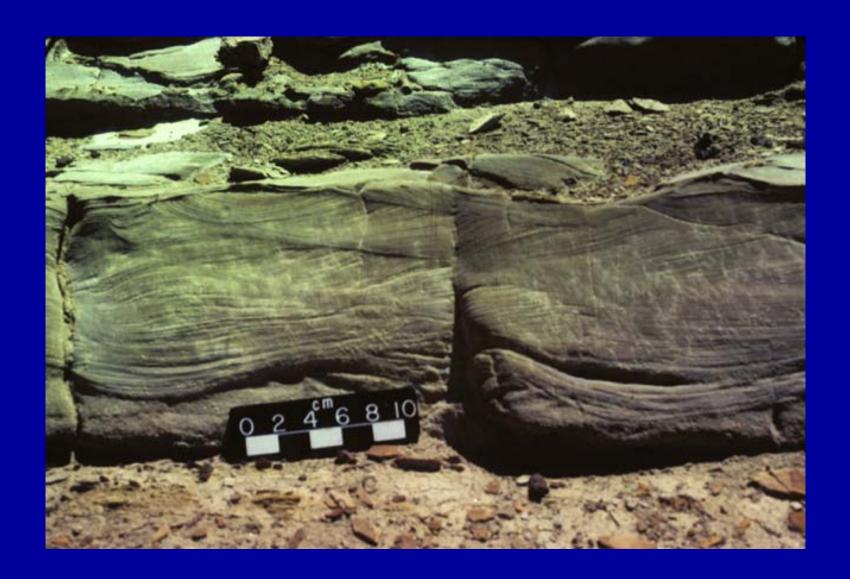
(Busby-Spera and Boles, 1986)

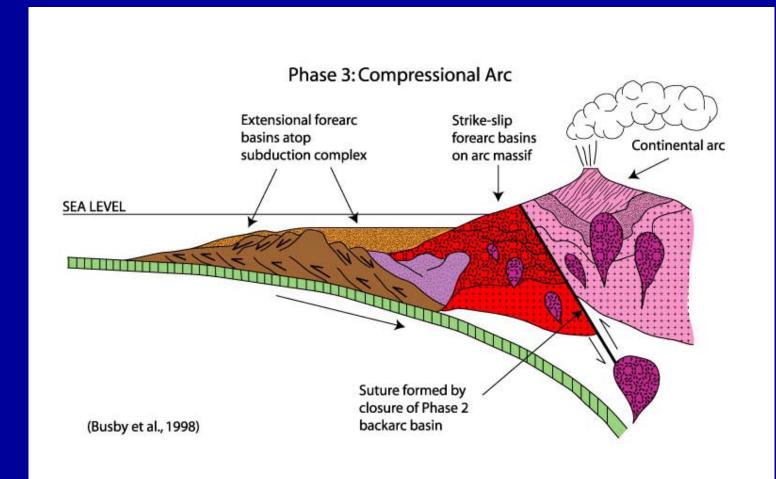


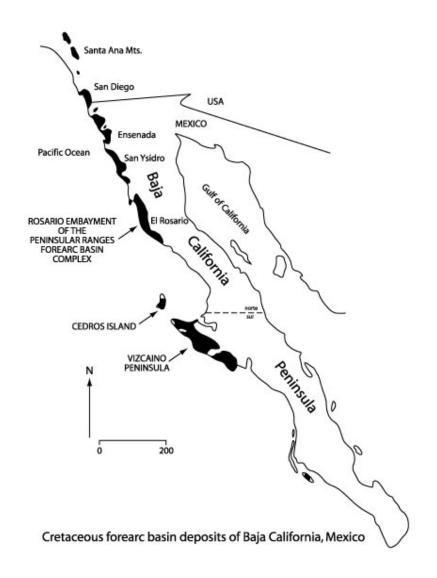




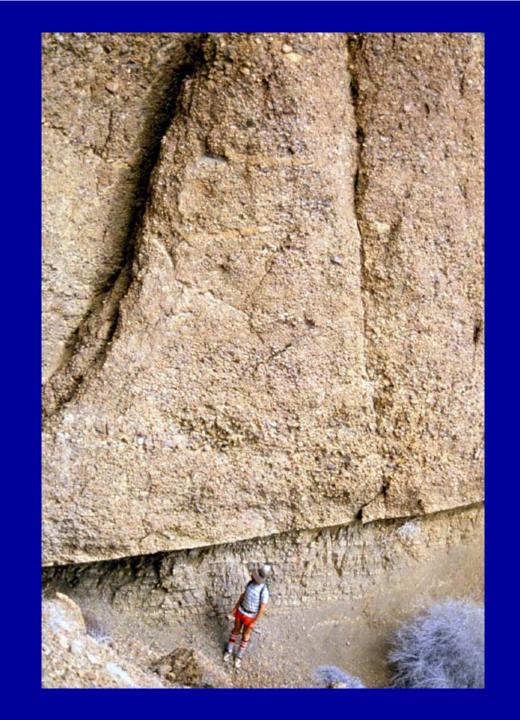


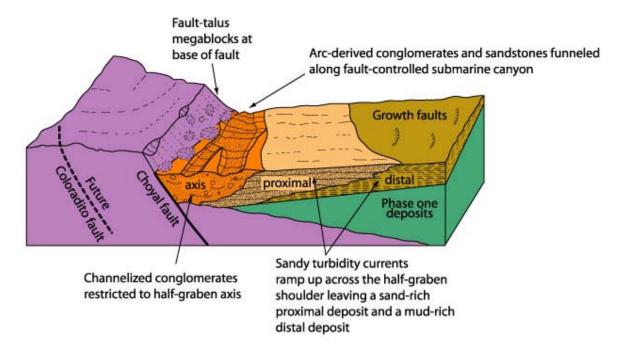




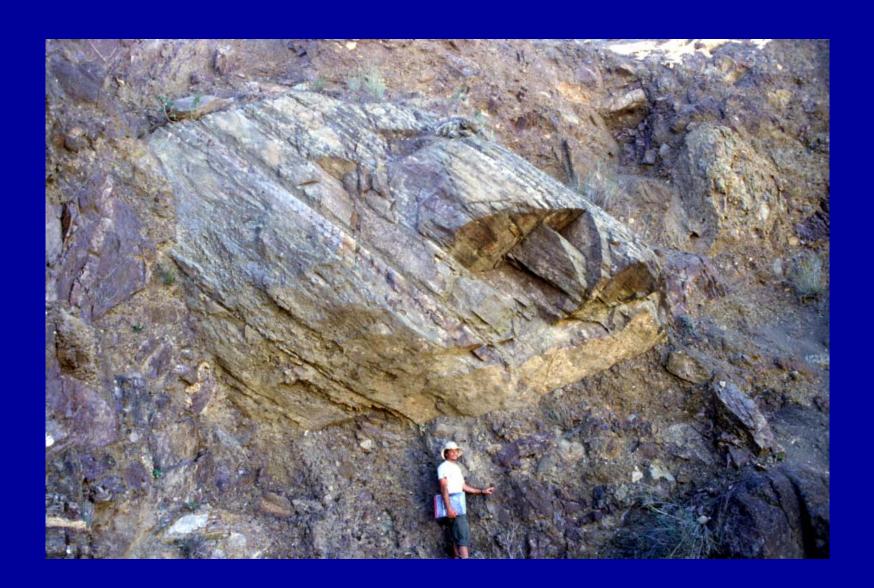


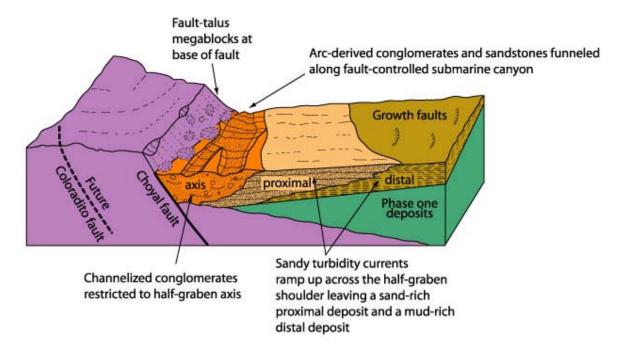






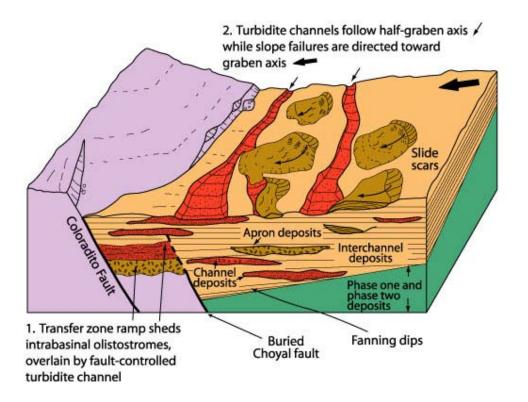
I. EXTENSIONAL FOREARC BASIN ATOP SUBDUCTION COMPLEX





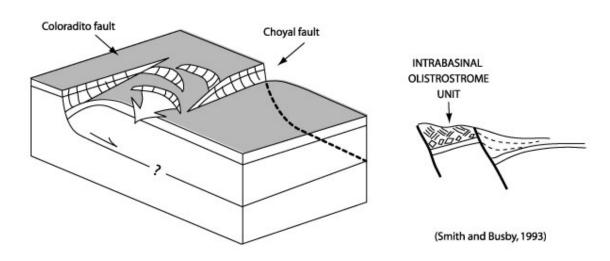
I. EXTENSIONAL FOREARC BASIN ATOP SUBDUCTION COMPLEX





II. EXTENSIONAL FOREARC BASIN ATOP SUBDUCTION COMPLEX

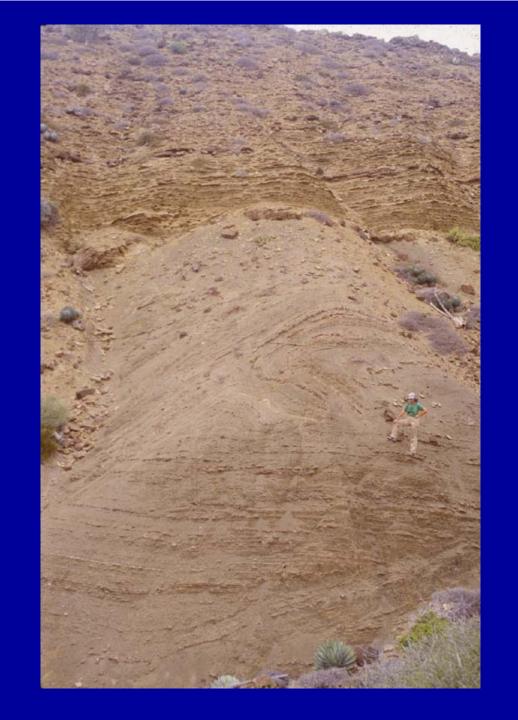
Growth of Coloradito fault *en echelon* with Choyal fault produces transfer zone ramp that sheds olistrostrome

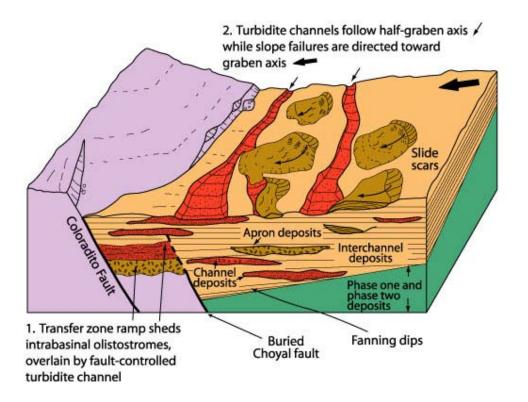




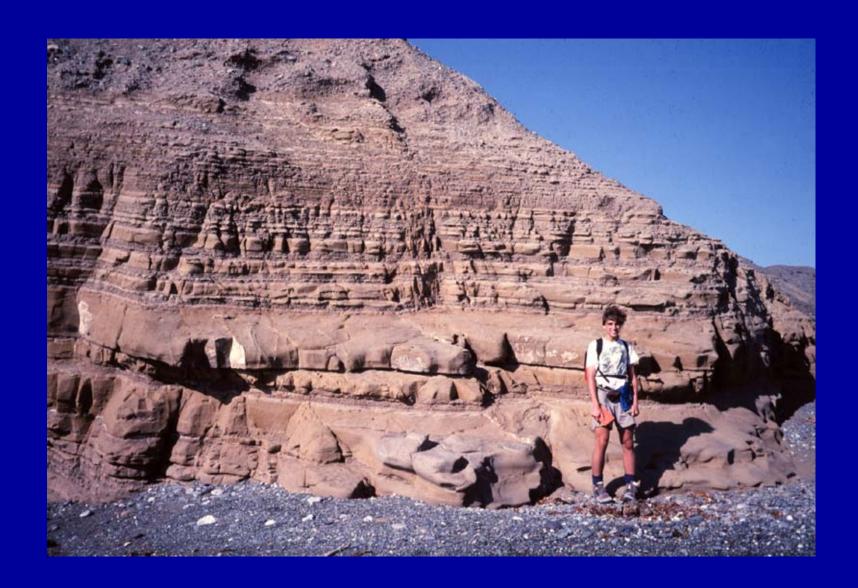


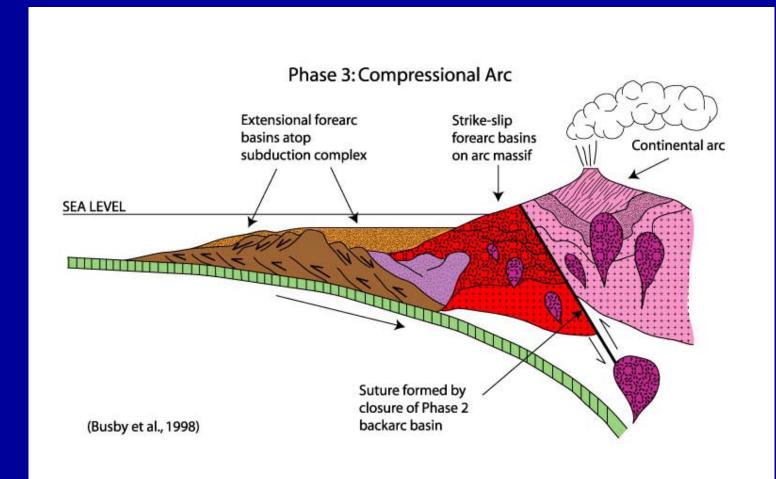




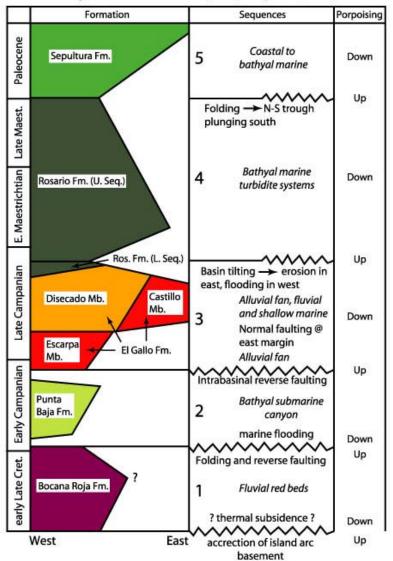


II. EXTENSIONAL FOREARC BASIN ATOP SUBDUCTION COMPLEX

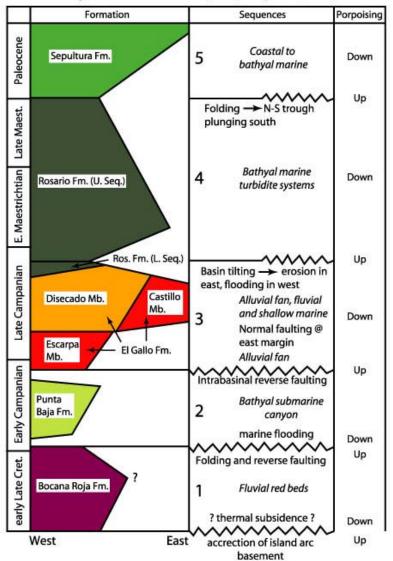


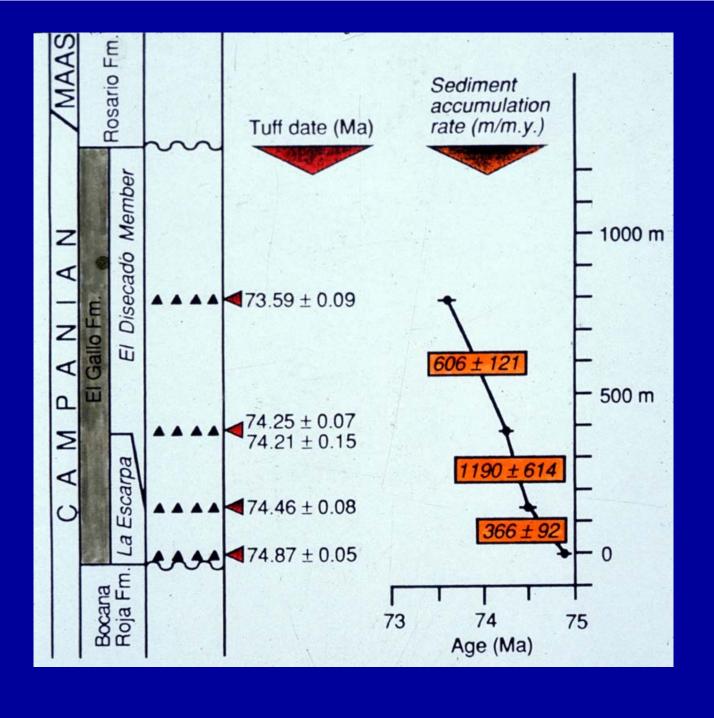


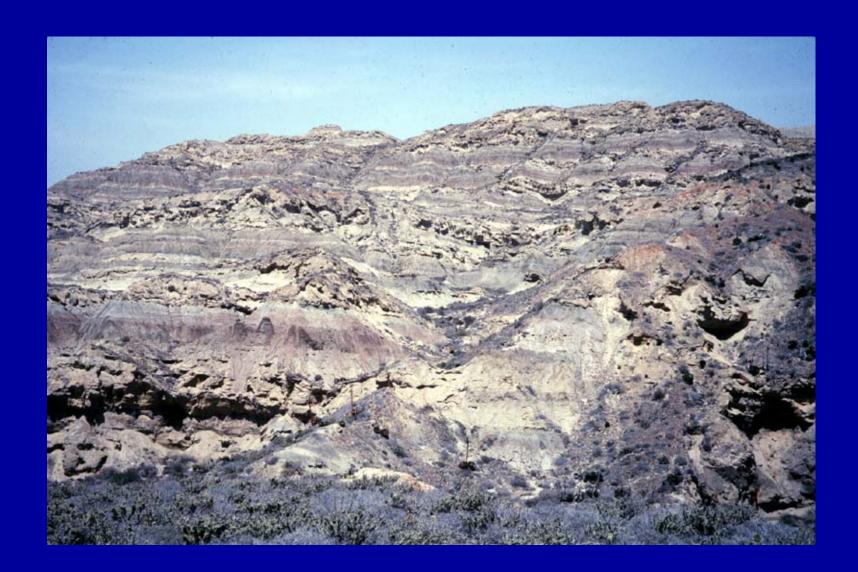
Sequence Stratigraphy of the Rosario Embayment of the Peninsular Ranges Forearc Basin Complex (Busby et al., 1998)

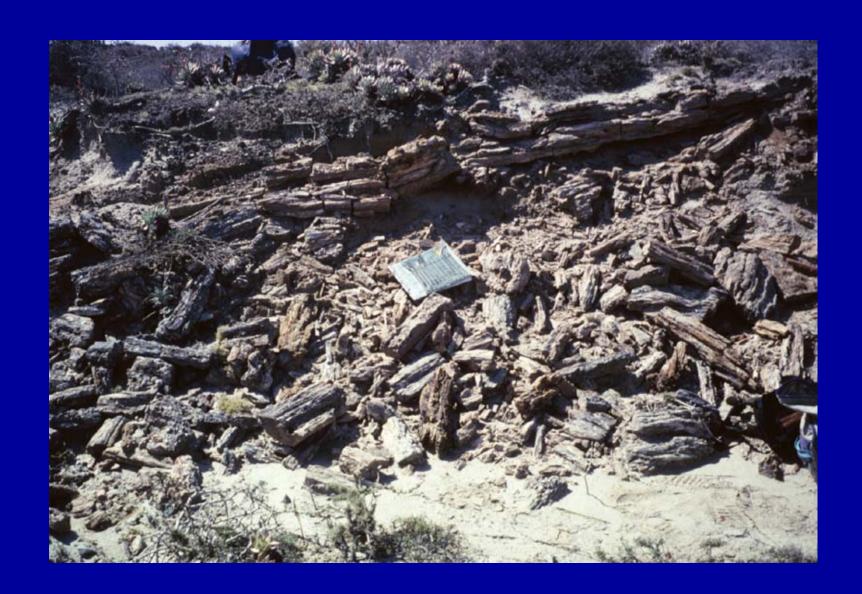


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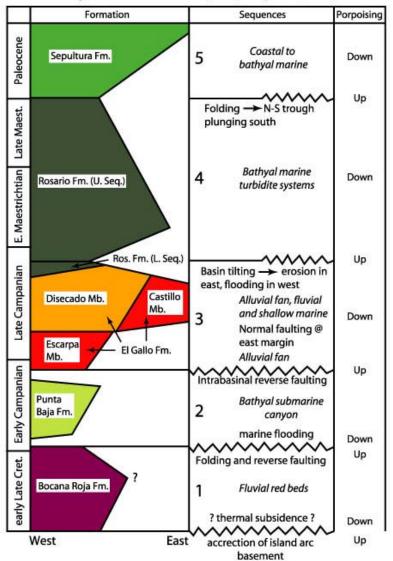




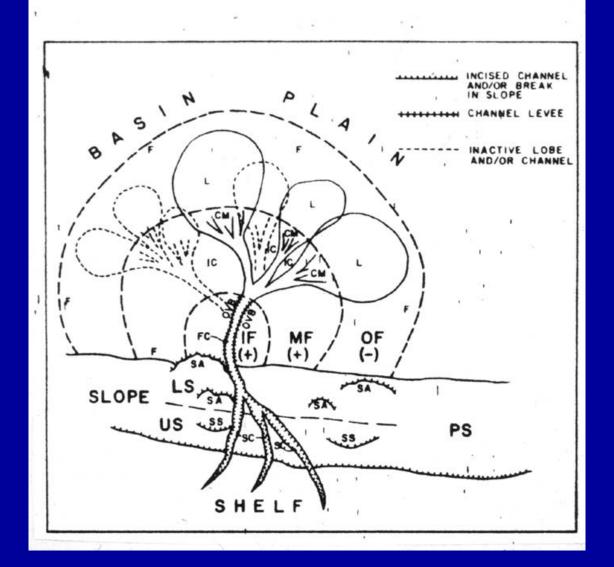




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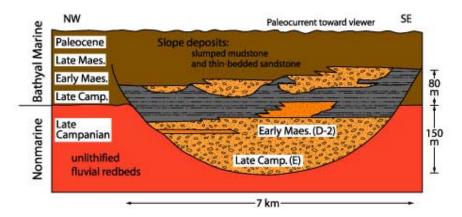
Ingersoll, 1979



San Carlos submarine canyon, Rosario Group, Peninsular Ranges forearc basin complex

(Morris and Busby-Spera, 1988)

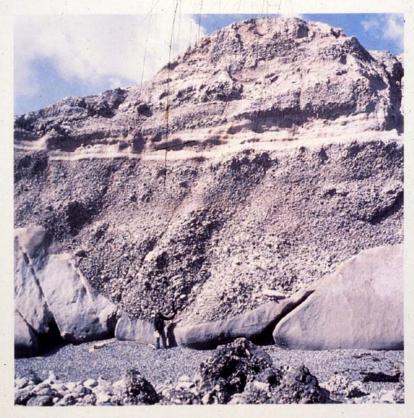
- Erosively-based feature filled wth bathyal marine deposits
- Canyon fill bound laterally by contemporaneous slope deposits



- Unlithified redbeds downdropped >1 km and eroded by turbidity currents
- Lower conglomerate sandstone unit: amalgamated channels, intraformational slide blocks 100m long, and giant load structures
- Middle mudstone sandstone unit: aggrading turbidite channel with traction structures; slumps in mudstones indicate high axial gradients
- Upper conglomerate sandstone unit: more isolated channels, migrate S.E. with time

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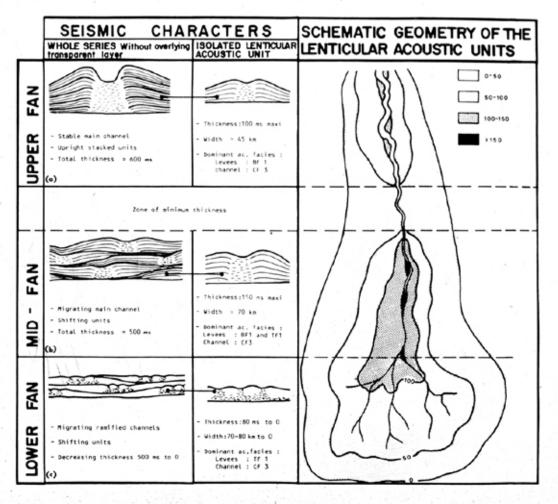
A·A·P·G BULLETIN







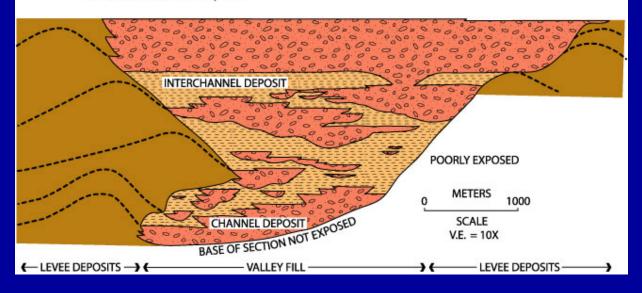
Droz and Bellaiche, 1985



Droz and Bellaiche, 1985 Active main channel Profile OW ENE WSW s Old abandoned channel New active channel Profile C ENE WSW



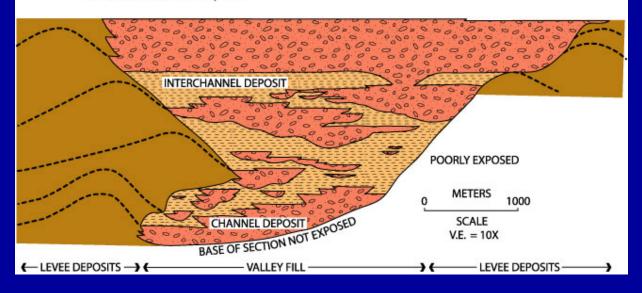
Aggradation of valley confined by levees leads to vertical stacking of channel (and interchannel) deposits

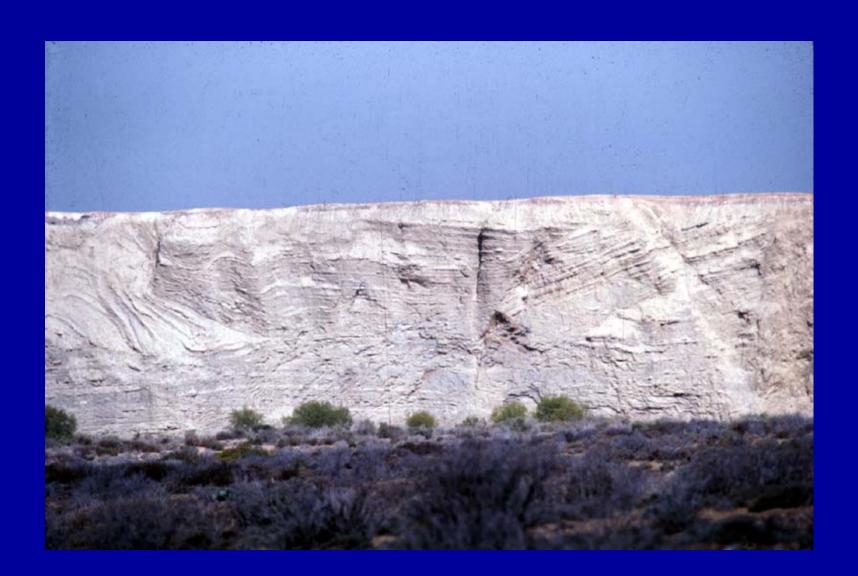




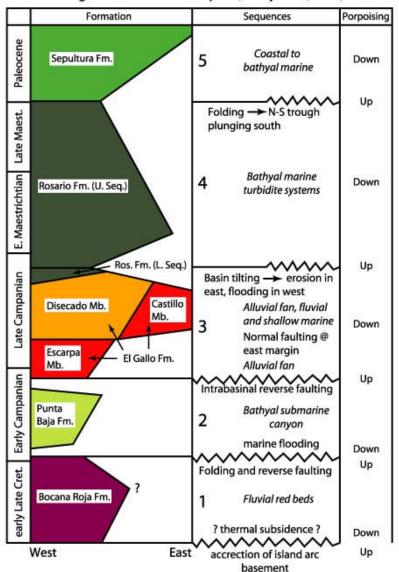


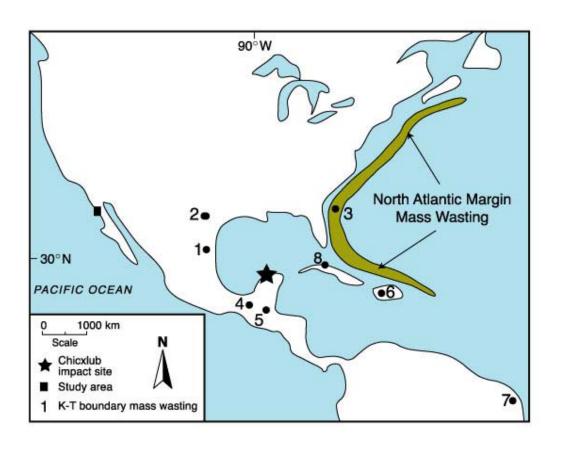
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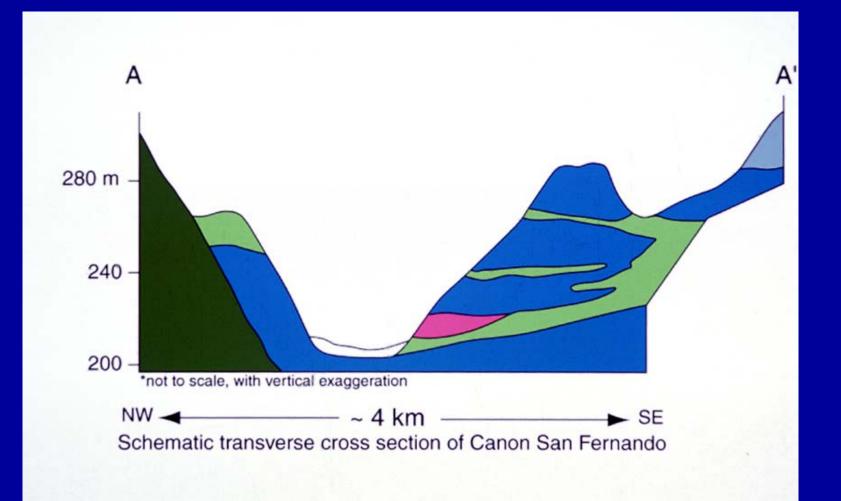


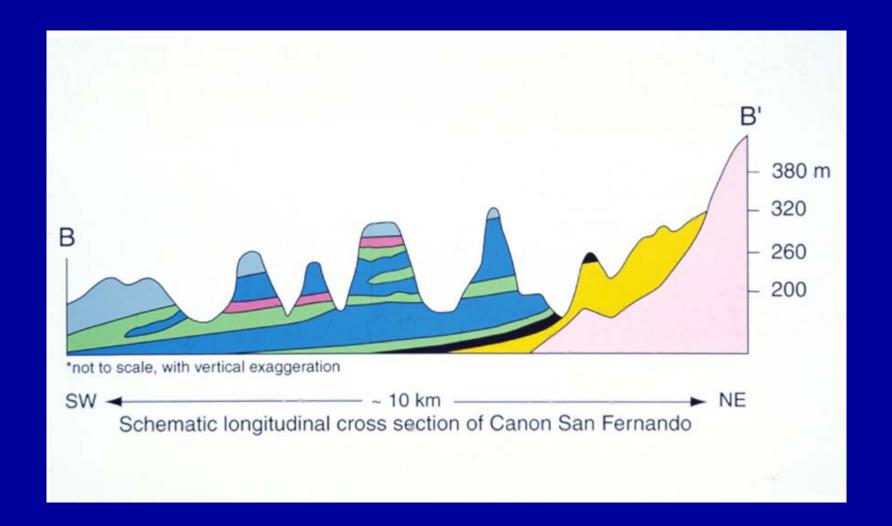


Sequence Stratigraphy of the Rosario Embayment of the Peninsular Ranges Forearc Basin Complex (Busby et al., 1998)

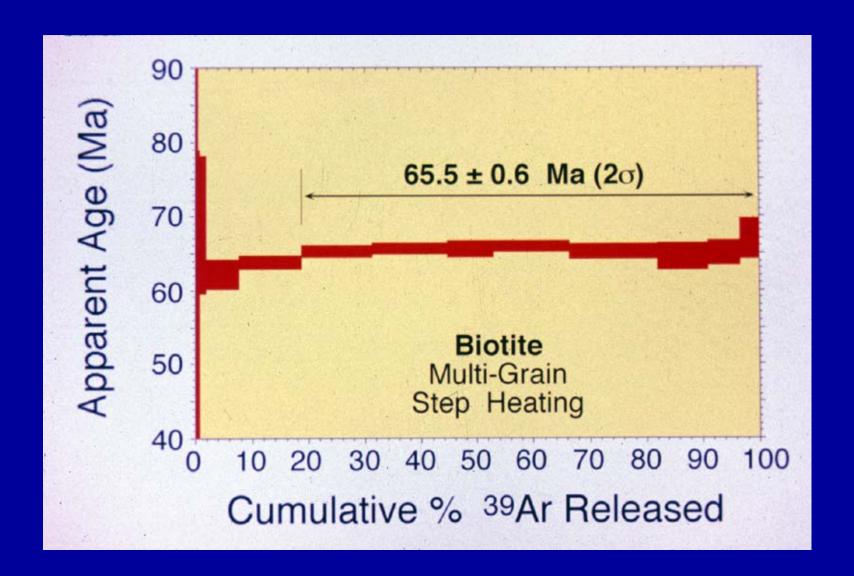




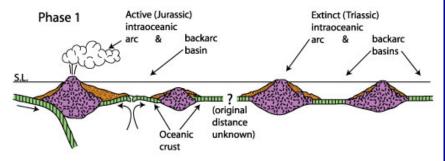








Strongly Extensional Arc



Moderately Extensional Arc

