# Quantifying sediment supply to continental margins: Application to the Paleogene Wilcox Group, Gulf of Mexico 

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## ESTIMATION OF SEDIMENT LOAD FOR THE PALEOGENE WILCOX GROUP USING SHELF-MARGIN CLINOFORMS

## Introduction and Methodology

Petter et al. (2013) proposed an inversion scheme to estimate the sediment flux from the dimensions of shelf-margin clinoforms and the progradation and aggradation rates of the related shelf edge. This approach is able to give a reliable estimate of sediment flux with several limitations (e.g., influences from mass transport complexes and canyons; see also Salazar, 2014).

The sediment flux of the entire shelf-margin clinoform can be estimated by equation S :

$$
\begin{equation*}
q_{s}(x)=\epsilon_{\mathrm{bed}}\left[P \eta(\mathrm{x})+\int_{x}^{L} A(x) d x\right] \tag{S1}
\end{equation*}
$$

where $q_{s}(x)$ is the sediment flux of the shelf-margin clinoform at the down-basin coordinate $x ; \mathrm{\epsilon}_{\text {bed }}$ is the 1-porosity ( $20 \%$ is used in this study); $P$ is the shelfmargin progradation rate; $\eta(\mathrm{x})$ is the clinoform elevation at $x ; A$ is the shelf-margin aggradation rate; and $L$ is the downdip location of the distal clinothem pinch-out.

The two-dimensional (2-D) estimate is extended into the three-dimensional estimate by summing up the results of each clinoform along the Wilcox shelf margin within each catchment (equation S2). The estimated sediment flux (units of length/time) is then transferred into sediment load (units of mass/time) by multiplying by an uncompacted grain density of $2 \mathrm{t} / \mathrm{m}^{3}$ :

$$
\begin{equation*}
\mathrm{Q}_{s}=\int_{0}^{L} q s d l \tag{S2}
\end{equation*}
$$

where $\mathrm{Q}_{s}$ is the total sediment flux along one shelf margin; $L$ is the length of the shelf margin; $q_{s}$ is the sediment flux of each shelf-margin clinoform.

## Explanation of Inputs

An interpreted 2-D seismic line for onshore-offshore Gulf of Mexico in Peel et al. (1995) is selected to estimate the dimensions of Wilcox shelf-margin clinoforms (Figure S1). The shelf-margin clinoform is divided into three parts: topset (shelf deposits), foreset (slope deposits), and bottomset (basin floor deposits) based on its gradient and thickness variations. The shelf edge between topset and foreset is recognized as the rollover from topset ( $\sim 1.2^{\circ}$ gradient, estimated from the original seismic line) to foreset ( $\sim 3.5^{\circ}$ gradient, estimated from the original seismic line). The shelf edge is located at $77 \mathrm{~km}(48 \mathrm{mi})$ from the left boundary of the cross section. From 77 to 235 km ( 48 to 146 mi ), the slope-basin floor deposits thin basinward following an exponential trend. The toe of slope is difficult to define because of interruption by faults and salt tectonics. A previous study suggested the slope length is no longer than $270 \mathrm{~km}(168 \mathrm{mi})$ based on the modern graded slope angle ( $\sim 0.8^{\circ}$ ) in the Gulf of Mexico (McDonnell et al., 2008). Our estimate shows a steeper slope angle and suggests a shorter slope length. The slope is defined by the thickness of accumulated strata. The toe of the slope is approximately located at the area where thickness tends to decay slowly or be nearly constant. The thickness of the Wilcox interval enters a less than 2000-m (6562$\mathrm{ft})$ zone from $160 \mathrm{~km}(99 \mathrm{mi})$ with slower decay rate and becomes even more stable from 235 km ( 146 mi ) $\left(\sim 1300 \mathrm{~m}[\sim 4265 \mathrm{ft}]\right.$ thick and $\sim 8.5 \times 10^{-5} \mathrm{~m} / \mathrm{yr}[\sim 2.8 \times$ $\left.10^{-4} \mathrm{ft} / \mathrm{yr}\right]$ sedimentation rate on average). We, therefore, suggest the slope length is close to $84 \mathrm{~km}(52 \mathrm{mi})$ and at least below $159 \mathrm{~km}(99 \mathrm{mi})$, which is close to the slope length in the Gulf of Mexico paleogeography map created by Fulthorpe et al. (2014). Pinch-out position and depth of the clinoform are estimated based on the exponential spatial decay of clinoform thickness as suggested in Petter et al. (2013) (equation S3).


Figure S1. (A) Gulf of Mexico depositional-dip seismic cross section, orange-colored interval is Wilcox deposits (after Peel et al., 1995). (B) Topography of Top Wilcox and thickness variation of Wilcox Group through depositional-dip based on panel A; note the transition from exponential thinning ( $77-235 \mathrm{~km}$ [48-146 mi]) to nonexponential thinning interval ( $>235 \mathrm{~km}$ [ $>146 \mathrm{mi}$ ]). (C) Sketch showing Wilcox shelfmargin clinoform geometry. $R^{2}=$ coefficient of determination.

$$
\begin{equation*}
x_{0}=\frac{\ln \left(T_{p} / a\right)}{k} \tag{S3}
\end{equation*}
$$

$x_{0}$ is the distal pinch-out distance; $T_{p}$ is the pinch-out thickness ( 10 m [33 ft] is used in this study); $a$ is the initial coefficient; and $k$ is the decay constant, determined by the slope geometry.

The clinoform height is estimated as 8700 m $(28,500 \mathrm{ft})$, as the difference between shelf-edge
height and pinch-out height. The inputs of each variable are summarized in Figure S1 and Table S1.

The Wilcox shelf-margin map and the isopach map in previous publications (Galloway et al., 2000; Galloway, 2001) are used to estimate the shelfmargin progradation and aggradation rates. Each catchment is divided into several segments with the same shelf-margin progradation and aggradation rate. The progradation rate decreases from lower Wilcox to

Table S1. Summary of Input Parameters for Clinoform-Based Approach to Estimate Sediment Flux

| Parameters | Shelf-edge aggradation rate ( $A$ ) | Shelf-edge progradation rate $(P)$ | Pinch-out position, $X_{0}$ | Shelf-margin clinoform height, $\eta$ |
| :---: | :---: | :---: | :---: | :---: |
| Range | $\begin{aligned} & 3 \times 10^{-5} \text { to } 3 \times 10^{-4} \\ & \left(10^{-4} \times 10^{-3}\right)(\text { Table S2 }) \end{aligned}$ | 0-30 (0-19) (Table S2) | 713,900 (2,342,200) | 8700 (28,500) |
| Units | $\mathrm{m} / \mathrm{yr}$ ( $\mathrm{ft} / \mathrm{yr}$ ) | km/Ma (mi/Ma) | m (ft) | m (ft) |
| Source and approach | Galloway (2001) | Galloway et al. (2000) | Estimated from Peel et al. (1995) by equation S3 | Estimated from Peel et al. (1995) |

Table S2. Shelf-Margin Length, Shelf-Edge Progradation and Aggradation Rate, and Estimates of Sediment Load of Each Catchment

Table S2. Continued

| Catchment | Segment Number | Variables |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Shelf-Margin Length ( $L$ ), km (mi) |  |  |  |  |  | Shelf-Edge Progradation Rate (P), m/yr (ft/yr) |  |  |  |  |  | Shelf-Edge Aggradation Rate ( $A$ ), m/yr ( $\mathrm{f} / \mathrm{yr}$ ) |  |  |  |  |  | Sediment Load ( $Q_{s}$ ), Mt/yr |  |  |
|  |  | Comments |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  | Estimate from Galloway et al. (2011); consider 10\% error |  |  |  |  |  | Estimate from Galloway et al. (2000); if SE retrogradation, $P$ is assumed as 0 |  |  |  |  |  | Estimate from Galloway et al. (2000); consider 20\% error |  |  |  |  |  | Low | Mid | High |
|  |  |  |  |  |  |  |  |  |  |  |  |  | Rang |  |  |  |  |  |  |  |  |  |
|  |  |  | ow |  | Mid |  | igh |  | ow |  | Mid | Hi | igh |  |  | M | id |  | gh |  |  |  |
| Middle Wilcox |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Rio Grande | 1 | 41 | (25) | 45 | (28) | 50 | (31) | 0 | 0.000 | 0.0002 | (0.001) | 0.0004 | (0.001) | 7.2E-05 | (0.0002) | 9.0E-05 | (0.0003) | 1.1E-04 | (0.0004) | 4 | 6 | 8 |
|  | 2 | 204 | (127) | 227 | (141) | 250 | (155) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | 7.2E-05 | (0.0002) | $9.0 \mathrm{E}-05$ | (0.0003) | $1.1 \mathrm{E}-04$ | (0.0004) | 37 | 56 | 78 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 42 | 63 | 86 |
| Colorado | 1 | 230 | (143) | 256 | (159) | 282 | (175) | 0 | 0.000 | 0.0002 | (0.001) | 0.0004 | (0.001) | 1.4E-04 | (0.0005) | $1.8 \mathrm{E}-04$ | (0.0006) | 2.2E-04 | (0.0007) | 50 | 71 | 94 |
|  | 2 | 34 | (21) | 38 | (24) | 42 | (26) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | $1.4 \mathrm{E}-04$ | (0.0005) | $1.8 \mathrm{E}-04$ | (0.0006) | 2.2E-04 | (0.0007) | 10 | 15 | 20 |
|  | 3 | 29 | (18) | 32 | (20) | 35 | (22) | 0 | 0.000 | 0.0002 | (0.001) | 0.0004 | (0.001) | $1.6 \mathrm{E}-04$ | (0.0005) | $2.0 \mathrm{E}-04$ | (0.0007) | $2.4 \mathrm{E}-04$ | (0.0008) | 7 | 10 | 13 |
|  | 4 | 88 | (55) | 98 | (61) | 108 | (67) | 0 | 0.000 | 0 | 0.000 | 0 | 0.000 | $1.4 \mathrm{E}-04$ | (0.0005) | $1.8 \mathrm{E}-04$ | (0.0006) | $2.2 \mathrm{E}-04$ | (0.0007) | 19 | 27 | 35 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 86 | 122 | 162 |
| Mississippi | 1 | 268 | (167) | 298 | (185) | 328 | (204) | 0 | 0.000 | 0 | 0.000 | 0 | 0.000 | 7.2E-05 | (0.0002) | 9.0E-05 | (0.0003) | 1.1E-04 | (0.0004) | 29 | 41 | 54 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 29 | 41 | 54 |
| Tennessee | 1 | 271 | (168) | 301 | (187) | 331 | (155) | 0 | 0.000 | 0 | 0.000 | 0 | 0.000 | 4.0E-05 | (0.0001) | 5.0E-05 | (0.0002) | 6.0E-05 | (0.0002) | 16 | 23 | 30 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 16 | 23 | 30 |
| Upper Wilcox |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Rio Grande | 1 | 191 | (119) | 212 | (132) | 233 | (145) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | 1.9E-04 | (0.0006) | 2.4E-04 | (0.0008) | 2.9E-04 | (0.0009) | 70 | 101 | 136 |
|  | 2 | 89 | (55) | 99 | (62) | 109 | (68) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | 1.9E-04 | (0.0006) | $2.4 \mathrm{E}-04$ | (0.0008) | 2.9E-04 | (0.0009) | 26 | 40 | 56 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 96 | 140 | 192 |
| Colorado | 1 | 29 | (18) | 32 | (20) | 35 | (22) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | 1.8E-04 | (0.0006) | 2.2E-04 | (0.0007) | 2.6E-04 | (0.0009) | 8 | 12 | 17 |
|  | 2 | 48 | (30) | 53 | (33) | 58 | (36) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | $1.4 \mathrm{E}-04$ | (0.0005) | $1.8 \mathrm{E}-04$ | (0.0006) | 2.2E-04 | (0.0007) | 14 | 20 | 28 |
|  | 3 | 146 | (91) | 162 | (101) | 178 | (111) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | $9.6 \mathrm{E}-05$ | (0.0003) | $1.2 \mathrm{E}-04$ | (0.0004) | 1.4E-04 | (0.0005) | 21 | 35 | 52 |
|  | 4 | 16 | (10) | 18 | (11) | 20 | (12) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | 5.6E-05 | (0.0002) | 7.0E-05 | (0.0002) | 8.4E-05 | (0.0003) | 3 | 4 | 5 |
|  | 5 | 59 | (36) | 65 | (40) | 72 | (44) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | 5.6E-05 | (0.0002) | 7.0E-05 | (0.0002) | 8.4E-05 | (0.0003) | 5 | 9 | 14 |
|  | 6 | 27 | (17) | 30 | (19) | 33 | (21) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | 5.6E-05 | (0.0002) | 7.0E-05 | (0.0002) | 8.4E-05 | (0.0003) | 4 | 7 | 9 |
|  | 7 | 81 | (50) | 90 | (56) | 99 | (62) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | $7.2 \mathrm{E}-05$ | (0.0002) | $9.0 \mathrm{E}-05$ | (0.0003) | $1.1 \mathrm{E}-04$ | (0.0004) | 9 | 16 | 24 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 63 | 103 | 149 |

Table S2. Continued

| Catchment | Segment Number | Variables |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | Shelf-Margin Length (L), km (mi) |  |  |  |  |  | Shelf-Edge Progradation Rate (P), m/yr (ft/yr) |  |  |  |  |  | Shelf-Edge Aggradation Rate ( $A$ ), m/yr ( $\mathrm{ft} / \mathrm{yr}$ ) |  |  |  |  |  | Sediment Load ( $Q_{5}$ ), Mt/yr |  |  |
|  |  | Comments |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  | Estimate from Galloway et al. (2011); consider 10\% error |  |  |  |  |  | Estimate from Galloway et al. (2000); if SE retrogradation, $P$ is assumed as 0 |  |  |  |  |  | Estimate from Galloway et al. (2000); consider 20\% error |  |  |  |  |  | Low | Mid | High |
|  |  | Range |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  |  | Low |  | Mid |  | High |  | Low |  | Mid |  | High |  | Low |  | Mid |  | High |  |  |  |  |
| Mississippi | 1 | 95 | (59) | 105 | (65) | 116 | (72) | 0 | 0.000 | 0 | 0.000 | 0 | 0.000 | 3.2E-05 | (0.0001) | 4.0E-05 | (0.0001) | $4.8 \mathrm{E}-05$ | (0.0002) | 5 | 6 | 8 |
|  | 2 | 93 | (58) | 103 | (64) | 113 | (70) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | 3.2E-05 | (0.0001) | $4.0 \mathrm{E}-05$ | (0.0001) | $4.8 \mathrm{E}-05$ | (0.0002) | 4 | 10 | 17 |
|  | 3 | 25 | (16) | 28 | (17) | 31 | (19) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | 3.2E-05 | (0.0001) | $4.0 \mathrm{E}-05$ | (0.0001) | $4.8 \mathrm{E}-05$ | (0.0002) | 3 | 5 | 7 |
| Tennessee | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 12 | 21 | 32 |
|  | 1 | 42 | (26) | 47 | (29) | 52 | (32) | 0.004 | (0.013) | 0.006 | (0.020) | 0.008 | (0.026) | 3.2E-05 | (0.0001) | 4.0E-05 | (0.0001) | 4.8E-05 | (0.0002) | 5 | 8 | 11 |
|  | 2 |  | (76) |  | (85) |  | (93) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | $1.6 \mathrm{E}-05$ | (0.0001) | $2.0 \mathrm{E}-05$ | (0.0001) | $2.4 \mathrm{E}-05$ | (0.0001) | 3 | 9 | 17 |
|  | 3 |  | (88) |  | (98) |  | (108) | 0 | 0.000 | 0.002 | (0.007) | 0.004 | (0.013) | $1.6 \mathrm{E}-05$ | (0.0001) | $2.0 \mathrm{E}-05$ | (0.0001) | $2.4 \mathrm{E}-05$ | (0.0001) | 3 | 11 | 19 |
|  | Sum |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  | 12 | 28 | 47 |

[^0]middle and upper Wilcox, ranging from 0 to $30 \mathrm{~km} / \mathrm{Ma}$. The aggradation rate ranges from $3 \times 10^{-5}$ to $3 \times 10^{-4}$ $\mathrm{m} / \mathrm{yr}\left(10^{-4} \times 10^{-3}\right)$. The progradation and aggradation rate along the Wilcox shelf margin from west to east are shown in Table S2.

## Results

Following the approach described here, the estimated sediment load of each catchment for lower, middle, and upper Wilcox is plotted in Table S2. The results of lower and middle Wilcox are averaged to yield the sediment load of the Paleocene Wilcox Group when compared with BQART Monte Carlo simulation (BQART-MCS) results. We notice that the ranges for sediment load tend to be narrower in the clinoform method than in the BQART-MCS. This probably reflects that (1) we have used only one cross section to characterize the shelf-margin architecture, and it is likely that the dimensions of the shelf-margin clinoforms vary at different locations along depositional strike; and (2) aggradation and progradation rates are averages over millions of years, and probably do not capture the full range of progradation- and aggradationrate variability over shorter time scales.

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[^0]:    Each catchment is divided into several segments from west to east. Abbreviations: $\mathrm{Mt}=$ million tons; $\mathrm{SE}=$ shelf edge.

